1. Introduction

During the Frasnian, a 5000 km² carbonate platform developed in Belgium, showing environments ranging from restricted shallow-water lagoons and supratidal areas to a relatively deep external ramp with carbonate mounds. This carbonate platform is especially instructive because of a combination of extraordinary outcrops (“marble” quarries with large sawn sections) and a long history of paleontological study which has led to a refined stratigraphic framework (Boulvain et al., 1999). Carbonate mounds have been the subject of intense investigation carried out by generations of geologists (e.g. Lecompte, 1959; Tsien, 1975; Boulvain, 2007) but a few of these studies focused on the shallow-water part of the platform (Da Silva & Boulvain, 2004).

2. Geological setting

Southern Belgium belongs to the northern part of the Rhenohercynian fold and thrust belt. Frasnian carbonates and shales are exposed along the borders of the Dinant, Vesdre and Namur Synclinoria and in the Philippeville Anticline (Fig. 1). The platform can be divided in three main depositional areas characterized by a different facies association, carbonate production rate and sedimentary evolution.

During the Middle Frasnian, the most distal part of the platform (“southern belt”) is located along the southern border of the Dinant Synclinorium, the “intermediate belt” corresponds mainly to the Philippeville Anticline and the “northern belt” -the shallower- crops out in the northern part of the Dinant Synclinorium, the Namur Synclinorium and the Vesdre area.

2.1. The southern belt

The southern belt is characterized by carbonate mound sedimentation with associated flank and off-mound facies. Since the classical studies by Mailleux (1913), three levels of carbonate mounds were known. These are in ascending order the Arche, Lion and Petit-Mont Members, belonging respectively to the Moulin Liénaux, Grands Breux and Neuville Formations (Fig. 2). The famous Arche and Lion buildups are located in the vicinity of Frasnes, historical stratotype of the Frasnian. Recently, Boulvain et al. (2005) gave information about a set of outcrops located some distance from Frasnes: the La Boverie quarry, close to Rochefort, and the Moulin Bayot sections, close to Vodelée. At both locations, it was possible to study the whole middle Frasnian succession, starting near the base of the Arche Member and ending within the Lion Member. Moreover, at both locations, an additional buildup was recognized between the Arche and Lion Members. The presence of this additional buildup along all the south side of the Dinant Synclinorium is now supported by its occurrence in boreholes drilled in the Nord quarry at Frasnes. The name of La Boverie Member was introduced by Boulvain & Coen-Aubert (2006), as a subdivision of the Moulin Liénaux Formation, for the carbonate deposits lying between the Arche and Bieumont Members.

Figure 1. Geological map of southern Belgium with location of stops.
2.2. The intermediate belt

In the Philippeville Anticline, the carbonate mound-bearing levels were replaced by shales and argillaceous limestones (Pont de la Folle Formation) followed by bedded limestone consisting of open-marine facies and biostromes (Philippeville Formation).

2.3. The northern belt

Along the northern border of the Dinant Synclinorium ("northern belt"), the Middle Frasnian consists of bedded limestones, exhibiting a distinct proximal aspect with biostromes alternating with lagoonal facies followed by palaeosoils and lagoonal deposits.

During the Upper Frasnian, a general northern shift or retrogradation of the platform is observed: the southern belt extends into the Philippeville Anticline with spectacular development of Petit Mont Member mounds and the intermediate belt shifts to the Northern border of the Dinant Synclinorium. This belt is characterized by shales with two carbonate levels dominated by rugose corals or oncoids (Aisemont Formation) (Boulvain, 2001) (Fig. 2).

3. Facies and microfacies

Data comes from the detailed study of more than 5000 thin sections from 20 outcrops from the Dinant, Namur and Vesdre Synclinoria and from the Philippeville Anticline. In the following descriptions, microfacies are ordered from the most distal to the most proximal, or from the deeper to the shallower. However, this order is not always effective, due to lateral variations, especially in the more proximal parts of the platform. Microfacies are grouped in 4 main facies belts: carbonate mounds and flank deposits (M), external platform or ramp (E), biostromes (B) and internal platform (I) (Fig. 3). Tables 1 & 2 compile detailed sedimentological characteristics and bathymetrical interpretations for the different microfacies.

3.1. Carbonate atolls and mounds (M, southern and intermediate belts, Figs 4 & 5)

The analogy between closely related facies in stratigraphically distinct buildups was highlighted by Boulvain et al. (2001) who employed the same facies designation, i.e. a number following a specific letter for the member name (for example: A2 and L2) corresponding to Figs 4 to 7 (sedimentary models).
In this more synthetic fieldtrip guidebook, the facies numbers are simply preceded by “M” for “mound”. Eight facies were recognised in the buildups (Table 1), each characterized by a specific range of textures and assemblage of organisms (Boulvain, 2007):

<table>
<thead>
<tr>
<th>FACIES/ASSOCIATION</th>
<th>COLOR / TEXTURE / STRUCTURE</th>
<th>AUTOCHTHONOUS AND ALLOCHTONOUS BIOTA</th>
<th>PRESERVATION / TRANSPORT</th>
<th>ENERGY</th>
<th>INTERPRETATION</th>
<th>BATHYMETRY</th>
</tr>
</thead>
<tbody>
<tr>
<td>M1. Stromatactis, sponges</td>
<td>Red mudstone or wackestone</td>
<td>Sponges and iron bacteria</td>
<td>Preservation ↑ Transport ↓</td>
<td>Very low</td>
<td>Aphotic, below SWAZ</td>
<td>100-150m</td>
</tr>
<tr>
<td>M2. Stromatactis, corals and crinoids</td>
<td>Red or pinkish mudstone, floatstone</td>
<td>Sponges, corals, crinoids and iron bacteria</td>
<td>Preservation ↑ Transport ↓</td>
<td>Very low</td>
<td>Aphotic, below SWAZ</td>
<td>80-100m</td>
</tr>
<tr>
<td>M3. Stromatactis, corals and stromatoporoids</td>
<td>Grey, pinkish or greenish, floatstone, (rudstone)</td>
<td>Corals, crinoids, brachiopods, bryozoan and stromatoporoids</td>
<td>Preservation ↑ Transport ↓</td>
<td>Low, episodically moderate</td>
<td>Subphotic, close to SWAZ</td>
<td>60-80m</td>
</tr>
<tr>
<td>M4. Corals, peloids and dasycladales</td>
<td>Grey graineastone, rudstone</td>
<td>Corals, stromatoporoids, dasycladales, cyanobacteria</td>
<td>Preservation ~ Transport ~</td>
<td>Moderate</td>
<td>Euphotic, close to FWWAZ</td>
<td>30-60m</td>
</tr>
<tr>
<td>M5. Microbial limestone</td>
<td>Grey, bindstone baffletone</td>
<td>Corals, cyanobacteria and stromatoporoids</td>
<td>Preservation ↓ Transport ~</td>
<td>Moderate</td>
<td>Euphotic, close to FWWAZ</td>
<td>30-60m</td>
</tr>
<tr>
<td>M6. Dendroid stromatoporoids</td>
<td>Grey, rudstone, m-thick beds</td>
<td>Dendroid stromatoporoids and cyanobacteria</td>
<td>Preservation ↑ Transport ↓</td>
<td>High</td>
<td>In the FWWAZ</td>
<td>0-30m</td>
</tr>
<tr>
<td>M7. Loferites</td>
<td>Grey, laminar, graineastone-wackestone</td>
<td>Dendroid stromatoporoids and paleosiphonocladales</td>
<td>Preservation ~ Transport ↓</td>
<td>Low</td>
<td>Intertidal</td>
<td>0m</td>
</tr>
<tr>
<td>M8. Bioturbated limestone</td>
<td>Grey, dm-thick, wackestone-mudstone</td>
<td>Paleosiphonocladales and calcispheres</td>
<td>Preservation ↑ Transport ↓</td>
<td>Low</td>
<td>Subtidal</td>
<td>5-10m</td>
</tr>
</tbody>
</table>

**Lateral facies (M)**

<table>
<thead>
<tr>
<th>FACIES/ASSOCIATION</th>
<th>COLOR / TEXTURE / STRUCTURE</th>
<th>AUTOCHTHONOUS AND ALLOCHTONOUS BIOTA</th>
<th>PRESERVATION / TRANSPORT</th>
<th>ENERGY</th>
<th>INTERPRETATION</th>
<th>BATHYMETRY</th>
</tr>
</thead>
<tbody>
<tr>
<td>M9. Microbioclastic packstone</td>
<td>Dark grey, dm-thick, bedded packstone</td>
<td>Corals, brachiopods, ostracods, bryozoans</td>
<td>Preservation ↓ Transport ↑</td>
<td>Low</td>
<td>Below SWAZ</td>
<td>80-100m</td>
</tr>
<tr>
<td>M10. Bioclastic packstone, grainstone</td>
<td>Dark grey, dm-thick, bedded packstone-grainstone</td>
<td>Corals, brachiopods, ostracods, bryozoans and stromatoporoids</td>
<td>Preservation ↓ Transport ↑</td>
<td>Low, episodically moderate</td>
<td>Close to SWAZ</td>
<td>60-80m</td>
</tr>
<tr>
<td>M11. Peloids and intraclastic packstone and grainstone</td>
<td>Dark grey, dm-thick, bedded packstone-grainstone</td>
<td>Stromatoporoids corals, brachiopods and bryozoans</td>
<td>Preservation ↓ Transport ↑</td>
<td>Low, episodically moderate</td>
<td>Close to SWAZ</td>
<td>60-80m</td>
</tr>
</tbody>
</table>
Laterally to the buildup facies, thin-bedded bioclastic and intraclastic facies are observed, most elements of which underwent a certain transport. Frequent sorting and rounding of their elements characterize these facies. They are ordered according to their content and grain-size: microbioelastic, often argillaceous packstones with ostracodes, trilobites and criniconarids (M9); bioclastic packstones, grainstones and rudstones with intraclasts (M10) and packstones, grainstones and rudstones with peloids and intraclasts (M11).

Sedimentological evidence suggests that facies M1 and M2 correspond to iron bacteria-sponge-dominated laminar fenestral limestone, (M7); grey, bioturbated limestone (M8).

### Table 2. Main characteristics of the external, biostromal and internal facies.

<table>
<thead>
<tr>
<th>FACES</th>
<th>INTERPRETATION</th>
<th>ENERGY</th>
<th>PRESERVATION / TRANSPORT</th>
<th>AUTOCHTONOUS AND ALLOCOHTONOUS BIOTA</th>
<th>COLOR / TEXTURE / STRUCTURE</th>
<th>FACIES</th>
</tr>
</thead>
<tbody>
<tr>
<td>E1. Crinoidal packstone</td>
<td>Under SWAZ, external deposits</td>
<td>Low</td>
<td>Preservation ↓</td>
<td>Crinoids, ostracods</td>
<td>Dark grey dm beds, packstone to wackestone</td>
<td>E1. Crinoidal packstone</td>
</tr>
<tr>
<td>E2. Lithoclastic packstone to grainstone</td>
<td>Under SWAZ, slope deposits</td>
<td>Low</td>
<td>Preservation ↓</td>
<td>Clasts, crinoids, ostracods</td>
<td>Dark dm beds</td>
<td>E2. Lithoclastic packstone to grainstone</td>
</tr>
<tr>
<td>B1. Laminar stromatoporoids</td>
<td>Under or in SWAZ, biostromes</td>
<td>Episodic</td>
<td>Preservation ↓</td>
<td>Laminar stromatoporoids, ostracods and brachiopods</td>
<td>Light grey, plum. beds, covestone to rudstone</td>
<td>B1. Laminar stromatoporoids</td>
</tr>
<tr>
<td>B2. Low domical stromatoporoids</td>
<td>In SWAZ, biostromes</td>
<td>Mainly low, episodical agitation</td>
<td>Preservation ↓</td>
<td>Low domical stromatoporoids, crinoids</td>
<td>Grey, plum. beds, rudstone</td>
<td>B2. Low domical stromatoporoids</td>
</tr>
<tr>
<td>B3. Dendroid stromatoporoids</td>
<td>SWAZ</td>
<td>High</td>
<td>Preservation ↓</td>
<td>Dendroid stromatoporoids, ostracods, calcified matrix</td>
<td>Light grey, plum. beds, floatstone to bindstone to shell-fabric</td>
<td>B3. Dendroid stromatoporoids</td>
</tr>
<tr>
<td>I1. Amphipora, dasycladales and peloids</td>
<td>Subtidal, restricted</td>
<td>Low</td>
<td>Preservation ↓</td>
<td>Amphipora, dasycladales and peloids</td>
<td>Light grey, metric beds, floatstone to wackestone</td>
<td>I1. Amphipora, dasycladales and peloids</td>
</tr>
<tr>
<td>I2. Umbella packstone</td>
<td>Modest, low</td>
<td>Preservation ↓</td>
<td>Preservation ↓</td>
<td>Umbrella packstone, clasts, crinoids, ostracods, peloids</td>
<td>Grey, metric beds, heterogeneous, subnodular</td>
<td>I2. Umbrella packstone</td>
</tr>
<tr>
<td>I3. Mudstone</td>
<td>Subtidal, local, intertidal features</td>
<td>High</td>
<td>Preservation ↓</td>
<td>Decimetric to metric dark grey beds</td>
<td>Light grey, metric beds, floatstone to wackestone</td>
<td>I3. Mudstone</td>
</tr>
<tr>
<td>I4. Laminated limestone</td>
<td>Low to moderate</td>
<td>Preservation ↓</td>
<td>Preservation ↓</td>
<td>Laminated limestone</td>
<td>Grey dm beds, with undulated lamination</td>
<td>I4. Laminated limestone</td>
</tr>
<tr>
<td>I5. Paleosols</td>
<td>Low</td>
<td>Preservation ↓</td>
<td>Preservation ↓</td>
<td>Paleosols</td>
<td>Purplum beds, light grey, with pink staining</td>
<td>I5. Paleosols</td>
</tr>
</tbody>
</table>

Laminar fenestral limestone, (M7); grey, bioturbated limestone (M8).

Laminar fenestral limestone, (M7); grey, bioturbated limestone (M8).
communities, developing in a quiet aphotic and hypoxic environment (Bourque & Boulvain, 1993; Boulvain et al. 2001). M3 developed between storm wave base and fair weather wave base, in an oligophotic environment. Facies M5 developed close to fair weather wave base. Facies M6 and the fenestral limestone M7 correspond to an environment with slightly restricted water circulation. Facies M8 developed at subtidal depths in a quiet, lagoonal environment.

Microbioclastic packstones (M9) are characterized by an open-marine facies with brachiopods, bryozoans and crinoids, whereas bioclastic rudstones (M10) and intraclastic packstones or grainstones (M11) show a clear mound influence as most of the bioclastic and intraclastic material is derived from these buildups. (Humblet & Boulvain, 2001).

The main differences between the Middle and Late Frasnian mounds concern facies architecture and are a consequence of different palaeoceanographic settings. The large flattened Middle Frasnian Arche and Lion buildups show limited vertical differentiation, large-scale progradation features, extensive exportation of material towards off-reef environment and development of inner lagoonal facies. They grew offshore from a well-developed carbonate platform with a healthy carbonate factory. Middle Frasnian sea level fluctuations were relatively

![Figure 4](image4.png) Sedimentary model of Late Frasnian Petit-Mont Member mounds in the Philippeville Anticline.

![Figure 5](image5.png) Sedimentary models of Middle to Late Frasnian mounds along the Southern border of the Dinant Synclinorium.
3.2. Carbonate platform (E, B, I, intermediate and northern belts, Figs 6 & 7)

The ideal shallowing-upward facies succession starts with open-marine deposits corresponding to crinoidal packstones (E1). They are followed by biostromes with laminar stromatoporoids (B1), overturned and broken massive stromatoporoids (B2) and dendroid stromatoporoids (B3). Then, biostromes are overlain by subtidal lagoonal facies with *Amphipora*, paleosiphonocladales and peloids (I1), followed by mudstone (I3) and laminated peloidal facies (I4) in the intertidal zone. The subtidal and intertidal zones were cut by channels filled by *Umbella* and intraclasts (I2). The supratidal zone was characterized by paleosols (I5) (Da Silva & Boulvain, 2004) (Table 2).

An important sedimentological observation concerning platform evolution (intermediate and northern belts) is the apparent division seen in all the sections between an upper and a lower unit (Da Silva & Boulvain, 2002) (Figs 10 & 13). The lower unit is dominated in the intermediate belt by ramp facies with some biostromal interruptions, and in the northern belt by biostromes with lagoonal interruptions. The upper unit (lagoon) consists of an alternation of biostromes and lagoonal facies in the intermediate belt and of lagoonal facies (with paleosol) in the northern belt.

Within these sedimentological units, facies are stacked into metre-scale cycles, showing mainly shallowing-
Figure 8. Drowning of the Hautmont mound. Vodelée, Petit-Mont Member.

Figure 9. Logs of sections in the Hautmont quarry. Vodelée, Petit-Mont Member.
upward trends. Such cyclicity is common in Devonian shallow-water carbonates. Different kinds of cycles however, are identified here.

In the biostromal unit from the intermediate belt, sedimentation is mainly acyclic with the stacking of 10 cm-thick crinoidal beds, probably due to the deeper environment being less sensitive to minor relative sea-level variations.

In the lagoonal unit, the cycles are characterized by biostromes followed by lagoonal deposits and capped by intertidal laminites. In the northern belt, the biostromal unit shows one or few metres-thick cycles, with crinoid beds (the colonisation stage) followed by massive biostromes and lagoonal deposits and capped by intertidal laminites. The lagoonal unit is characterized by restricted subtidal and intertidal facies covered by, or transformed into paleosols. These cycles are not always complete.

4. Stops

4.1. Stop 1: Hautmont quarry, Vodelée, Petit-Mont Member

This stop is dedicated to a very spectacular Late Frasnian carbonate mound (Petit-Mont Member): the Hautmont mound near Vodelée. This active quarry is located on the SE end of the Philippeville Anticline. The central part of the mound is in nearly horizontal position. The top of the mound is well accessible and all stages of mound drowning are visible (Figs 8 & 9). The upper central part of the mound shows a core of grey microbial, coral stromatoporoid limestone (M5). This facies forms massive limestone with stylolites. Decimetre- to metre-scale growth cavities cemented by granular spar are abundant. Breccia is locally present. The fauna is dominated by subspherical coral colonies (Hankaxis, Phillipsastrea, Alveolites), Thamnopora, brachiopods and subordinate dendroid stromatoporoids (Amphipora). Renalcis is locally abundant. Thrombolitic structures and microbial mats are present. Within thrombolites, Renalcis is often associated with Palaeomicrocodium.

4.2. Stop 2: Beauchâteau quarry, Senzeilles, Petit-Mont Member

This abandoned marble quarry, located near the village of Senzeille in the SW part of the Philippeville Anticline, is the most spectacular outcrop of a Late Frasnian carbonate mound in Belgium. The mound is standing in subhorizontal position and large sawn sections expose facies ranging from the middle part of the mound (M3) to its top (M4 and 5). The upper central part of the mound shows interfingering between grey massive microbial facies and pink bedded bioclastic flank sediments. The left part of the quarry shows crinoid-rich argillaceous flank sediments.

Contradictory inferences about the initial mechanical state of carbonate mound mud appear to derive from field observations. The persistence of dips as high as 35° on the flanks of several mounds, the presence of lithoclasts in the grey limestone (M5) and the sharp distinct character of some fractures indicate early lithification. Conversely, plastic deformation of the sediment, presence of overturned coral colonies (very spectacular in the lower central panel of the quarry), formation of zebra structures by lateral compression, scarcity of hardgrounds and of sediment borings, and the irregular character of some synsedimentary fractures indicate an absence of early lithification. It appears that the sediment was initially sufficiently ductile to permit synsedimentary deformation, yet sufficiently coherent to have maintained open cavities (stromatactis) and significant relief. It is likely that the sediment had a gel-like consistency, probably related to the presence of significant quantities of organic matter.

Figure 10. Log of the Villers-le-Gambon section with microfacies and magnetic susceptibility.
4.3. Stop 3: Villers-le-Gambon section, Philippeville Formation

This section along a disused railway (Fig. 10) is located in the S of the Philippeville Anticline. It is one of the rare sections exposing almost completely the Philippeville Formation. The lower 50 m are dominated by dark argillaceous crinoidal limestones (E1) with some biostromal beds and chert intercalations. The upper 50 m are characterized by biostromal facies with well-developed massive stromatoporoids alternating with fenestral mudstones (I1) or branching stromatoporoids (I3) rudstones.

4.4. Stop 4: Tailfer quarry, Lustin Formation

Located along the Meuse river, the Tailfer section is part of the northern flank of a major anticline, close to the Northern border of the Dinant Synclinorium. The first part of the section starts along the main road with oolitic hematitic beds alternating with dark shales (Lower Frasnian, top of Presles Formation, Fig. 11). Carbonate production starts with crinoidal beds (boundary with the Lustin Formation) and development of biostromes. These biostromes are dominated by massive or laminar stromatoporoids. The succession continues in the Tailfer quarry with a splendid sawn section of a lamellar stromatoporoids biostrome (Fig. 12). This is followed by an alternation of lagoonal deposits and well-developed paleosols (Fig. 13).

Biostome edification by lamellar stromatoporoids (Facies B1) corresponds to the following ecological sequence (Fig.12): lamellar stromatoporoids rudstone with crinoids, brachiopods and packstone matrix characterizes the colonisation phase (B1/3, Fig.12B), and developed during relatively high energy events. Then, two other lamellar stromatoporoids facies with tabulate...
Figure 12. Lamellar stromatoporoids biostrome, Tailfer quarry, Lustin Formation.

Figure 14. MS evolution with sequence evolution, close up from Fig. 13.
corals (B1/2) and mud (B1/1), alternate in dm-scale units (Fig. 12C-D), probably in relation with short term higher energy events. Muddy microfacies settled in quiet to very quiet water (immediately below the storm wave zone) while reef bioclasts-rich microfacies corresponded to higher energy periods. Water energy however remained relatively weak (in comparison with microfacies B1/3), as indicated by good preservation of fossils.

5. Magnetic susceptibility and platform sediments

The trends in the MS signature are similar for all described stratigraphic sections. We will use the Tailfer section as a reference to describe the relationship between facies change and MS signature. Detailed results were published in Da Silva & Boulvain (2002, 2006, 2009a & b)

MS evolution seems to be related to different parameters. We will illustrate the relationship between MS and fourth (a), third (b) order sequences, with microfacies (c) and with the position of the section in the basin.

a) The first trend is a correlation between the cycles identified on the MS curve and the fourth order sequences. Each regressive trend corresponds to a MS peak on the MS evolution curve. On Fig. 14, the detail of some of these trends at the scale of fourth order sequences is presented and the link between MS and sequence evolution is obvious. Concerning the sequence 3 (Fig. 14), the first trend is a short transgressive event, with biostromal facies (facies B3) grading to external crinoidal deposits (facies E1). This transgressive phase corresponds to decreasing MS values. The second trend is an aggrading biostrome, mainly built by lamellar stromatoporoids (facies B1) that
Figure 17. MS curves and correlations for 4 sections from the carbonate platform of Belgium.
corresponds to low (around 0 m³/kg) and almost constant MS values. A regressive trend caps the sequence, with high domical stromatoporoids (facies B2) followed by laminar limestones (facies I4) from the intertidal zone; this regressive trend corresponds to increasing MS values. Within sequence 4, the transgressive phase is almost absent and is followed by an important aggradational phase (biostromes with laminar stromatoporoids, facies B1) and a regressive phase (high domical stromatoporoids, facies B2 followed by mudstone, facies 7). The corresponding MS evolution is almost constant during the aggrading phase and is followed by increasing values during the regressive phase. The sequences 5 and 10 (Fig. 14) also show transgressive trends followed by regressive trends with the MS curve respectively decreasing and increasing.

b) The second trend is a subdivision of the curve in two distinct parts (Fig. 13). The biostromal unit presents very low MS values (mean of 2x10⁻⁸ m³/kg). Just above the boundary between the two units, the values are increasing strongly. The upper portion of the curve is characterized by higher values (to means of 6.62x10⁻⁸ m³/kg), and corresponds to the lagoonal unit.

c) Fig. 15 presents the relationship between magnetic susceptibility and microfacies. On the horizontal axis, the microfacies succession is represented, with increasing proximality from the right to the left. The relationship between magnetic susceptibility and microfacies is obvious, with increasing MS related to proximality. Effectively, the external and biostromal microfacies show MS values around 2x10⁻⁸ m³/kg, while the lagoonal facies shows values around 6.7x10⁻⁸ m³/kg, with in details subtidal facies 5.5x10⁻⁸ m³/kg, intertidal facies 8.5x10⁻⁸ m³/kg and supratidal facies 5x10⁻⁸ m³/kg. It appears that paleosols always have lower MS values than intertidal deposits. This can be explained by the fact that these paleosols correspond mainly to subtidal deposits affected by pedogenesis. So the MS signal of the subtidal deposits seems to be preserved during emersion and pedogenesis. MS is related both to microfacies and to third order sequential evolution.

d) If we compare the MS values of the biostromal unit of the different sections (Fig. 16), the mean values are the highest in Barse which is the most proximal section and decrease distally. For the Villers section, the values are a little bit higher, maybe because of a lower sedimentation rate and local condensed intervals.

Correlations are made on basis of magnetic susceptibility peaks (events with the same pattern) which are considered isochronous (Crick et al., 1997). Third order to fourth order correlations are proposed on the basis of MS peaks (Fig. 17).

6. References


